

Observational constraints on the first indirect aerosol effect

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Acknowledgements

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Albedo parameterization
Indirect aerosol effect
Diurnal cycle of LWP
Diurnal cycle of CF

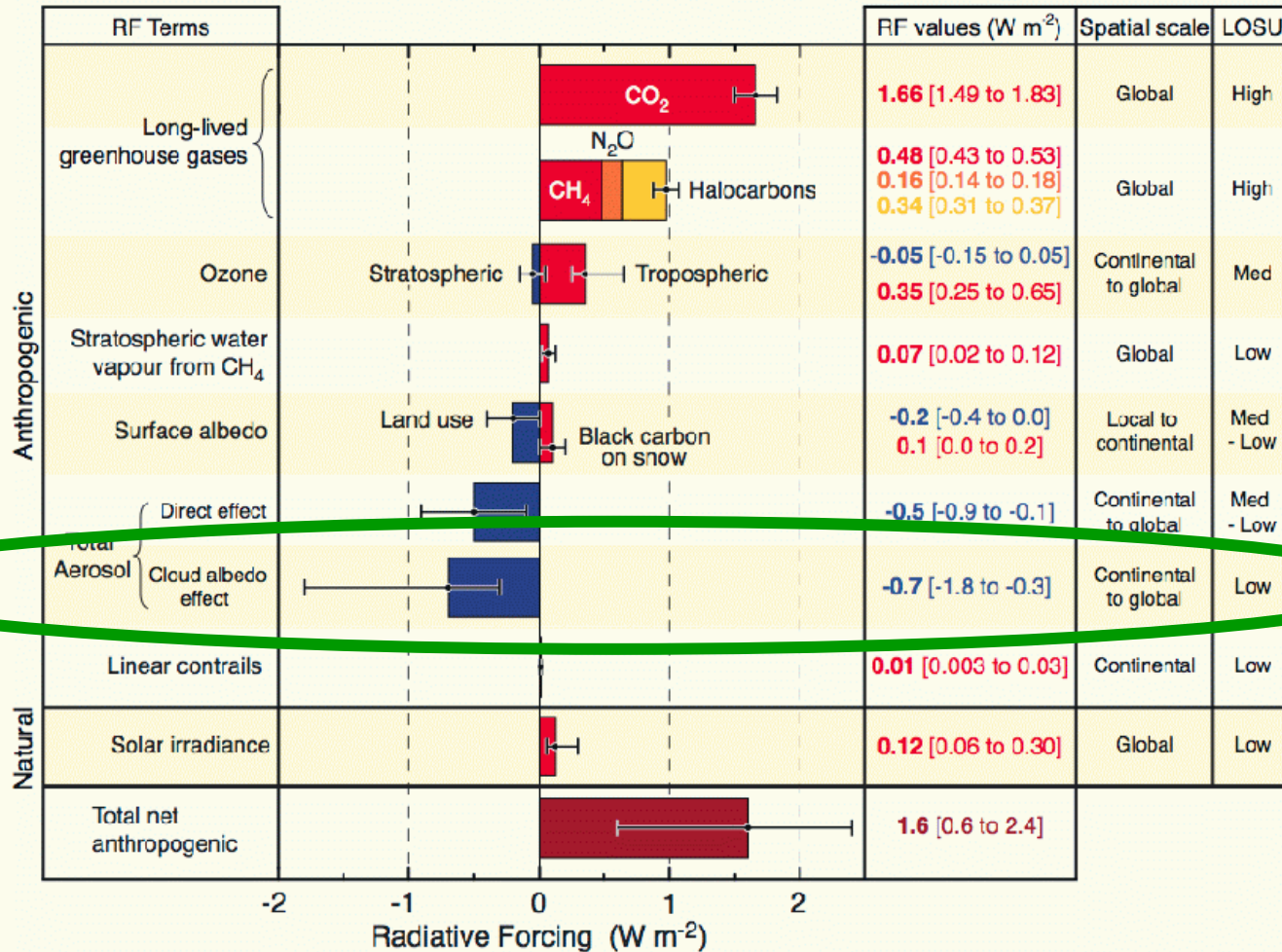
Brent Maddux
Lori Borg

UW-Madison
UW-Madison



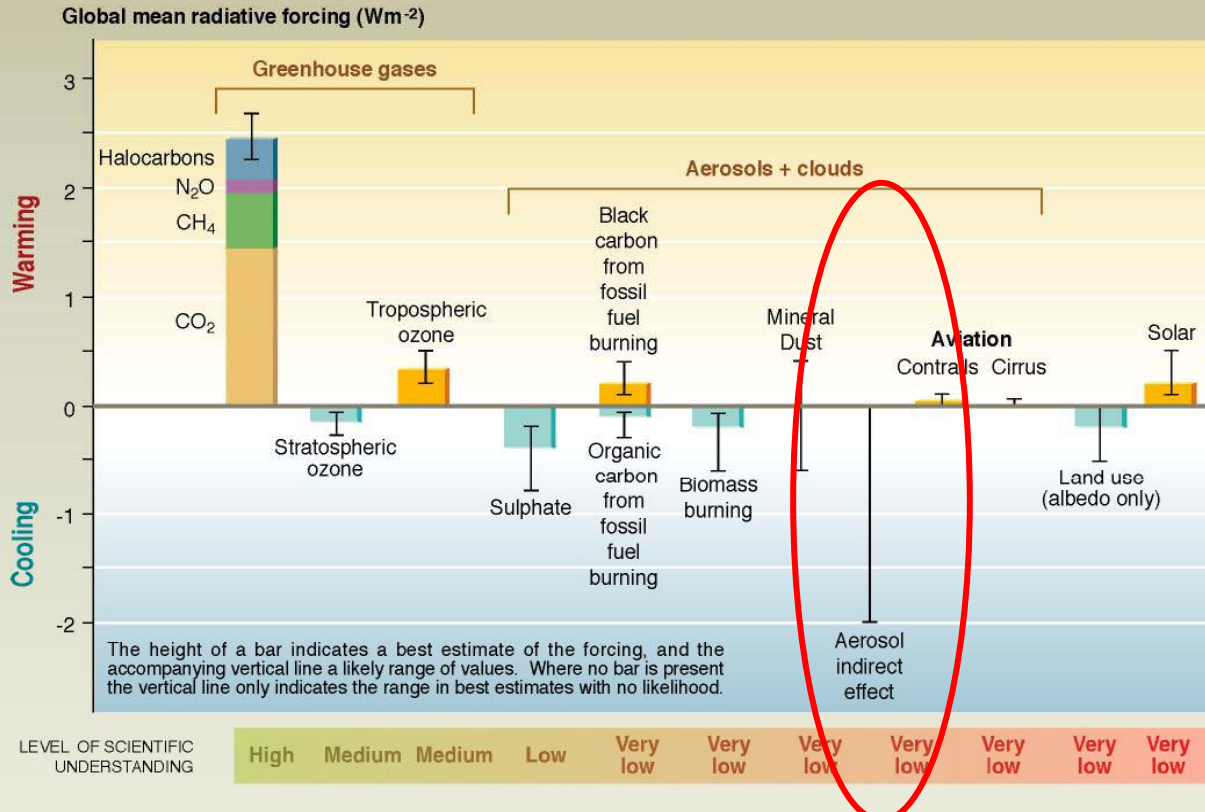
Data
Data

Radiative Forcing Components



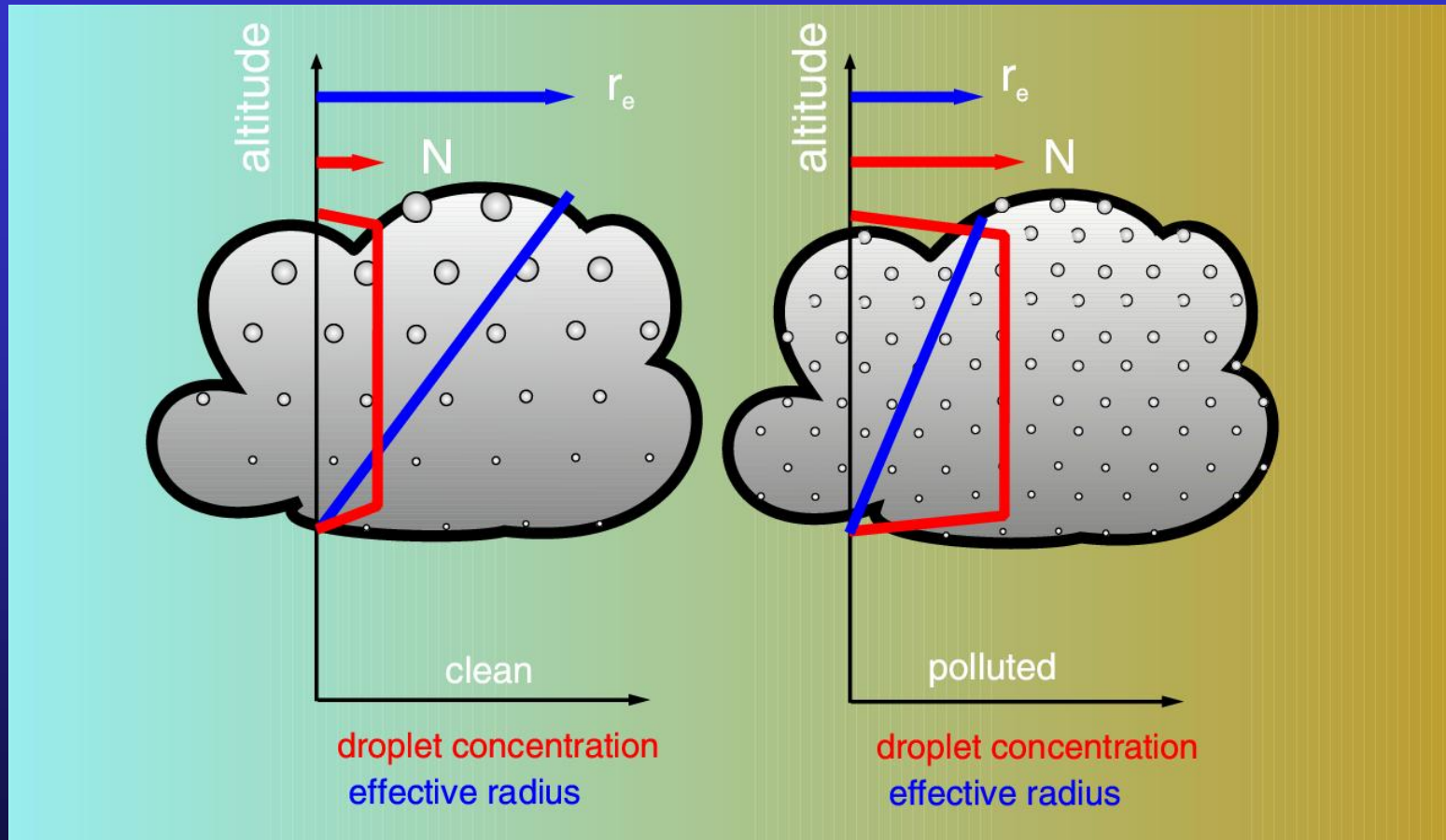
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Anthropogenic and natural forcing of the climate for the year 2000, relative to 1750



SYR - FIGURE 2-2

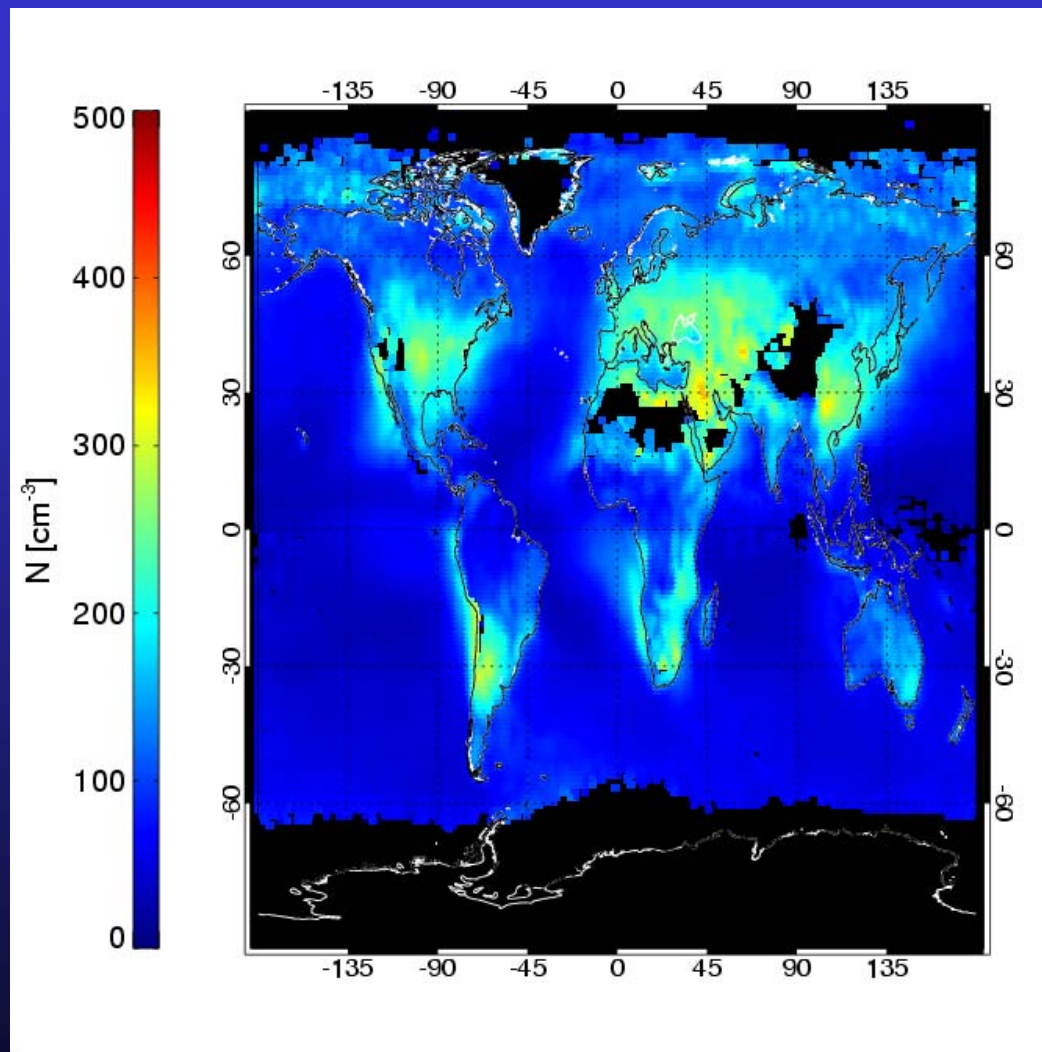
Microphysical changes due to pollution



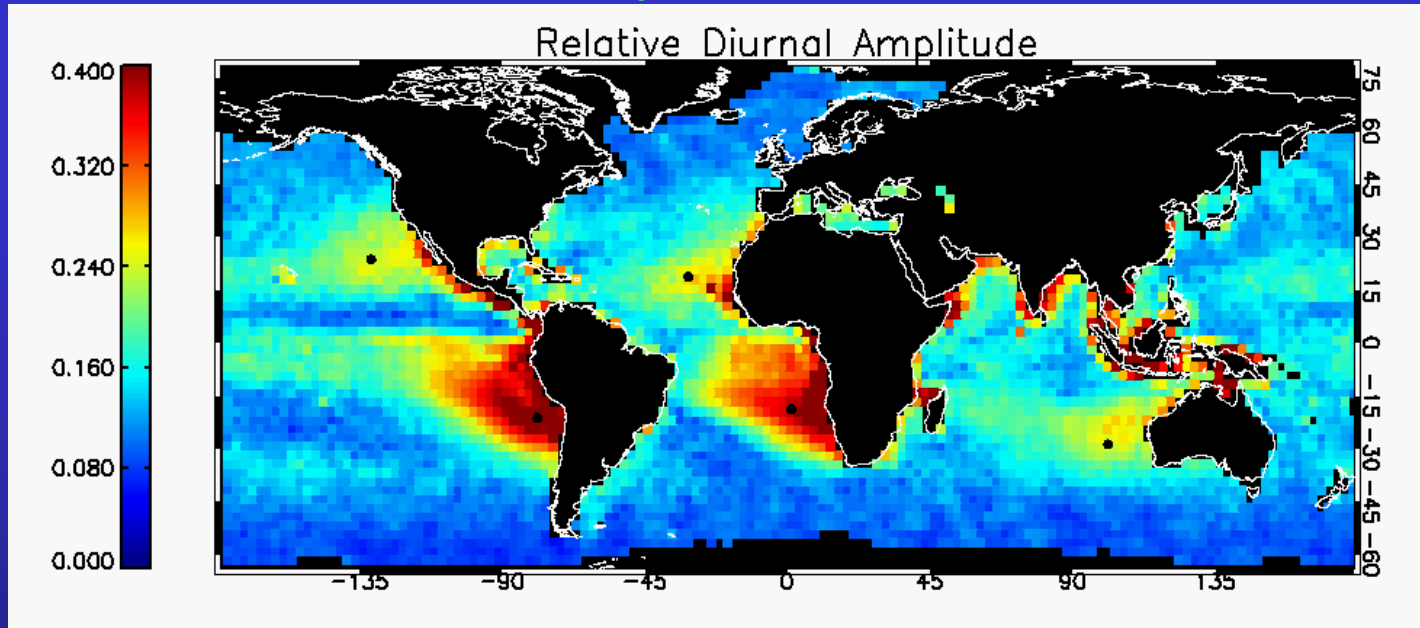
Data

- Cloud liquid water path diurnal cycle (O'Dell et al., 2007, J Climate, in press)
 - Cloud fraction diurnal cycle (ISCCP)
 - Cloud droplet number concentration (CDNC) (MODIS, Bennartz, 2007, JGR)
 - Aerosol & fine mode fraction (used as proxy for natural background fraction of CDNC)
 -
-

Cloud droplet number concentration from MODIS

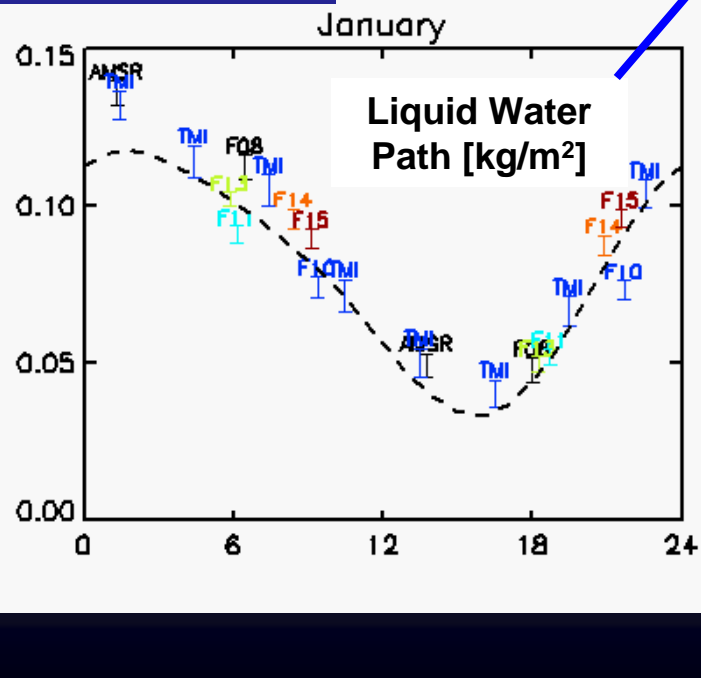
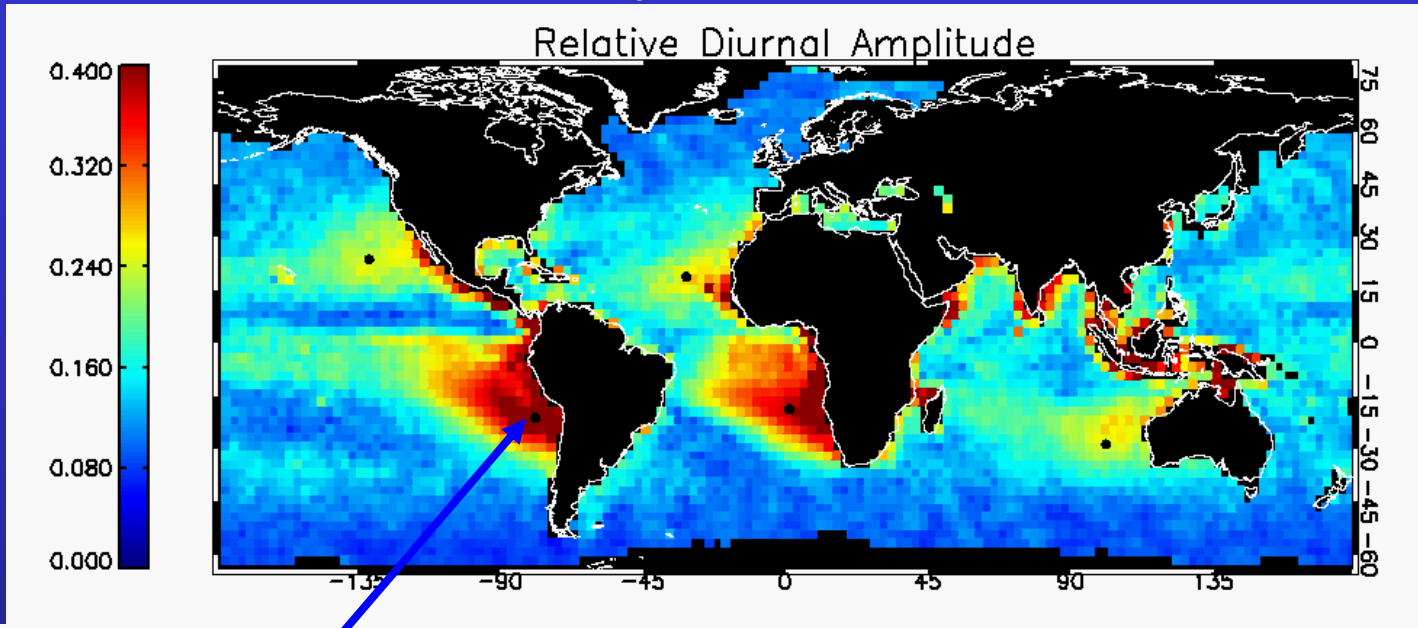


The diurnal cycle of clouds

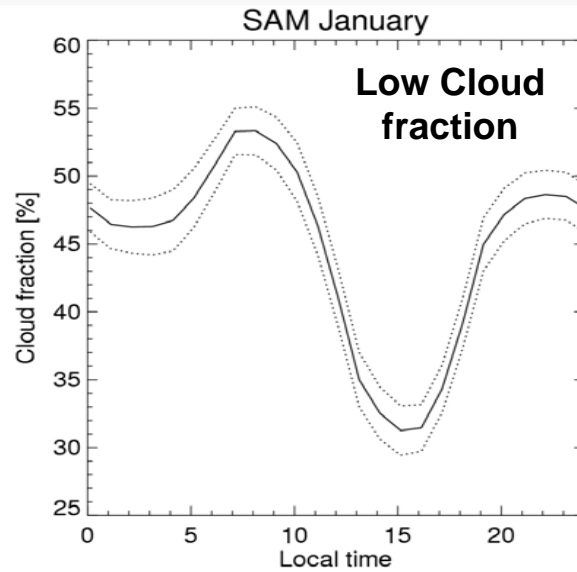
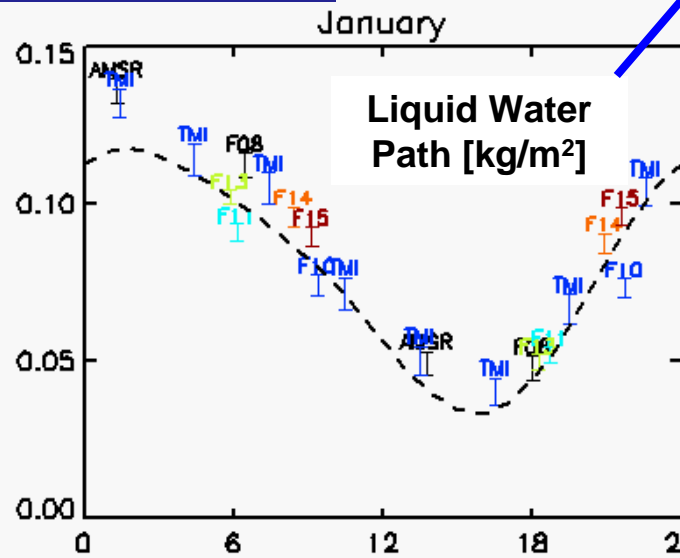
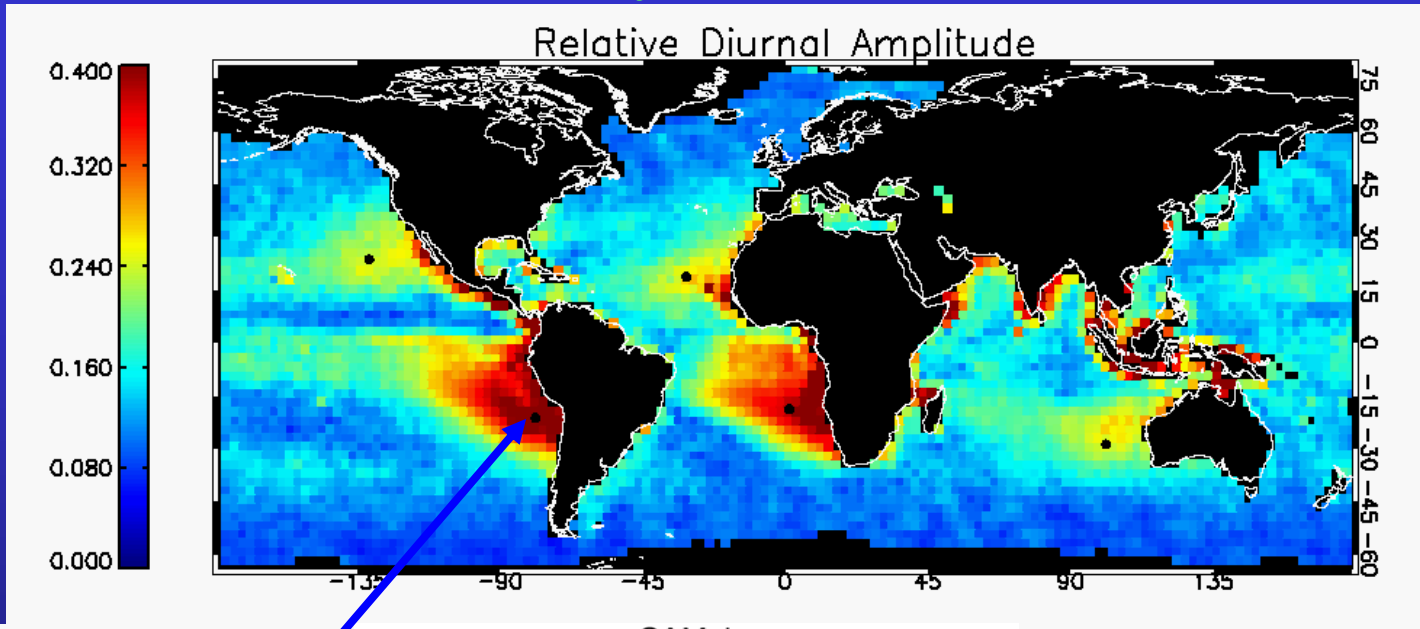


Local Time [hours]

The diurnal cycle of clouds

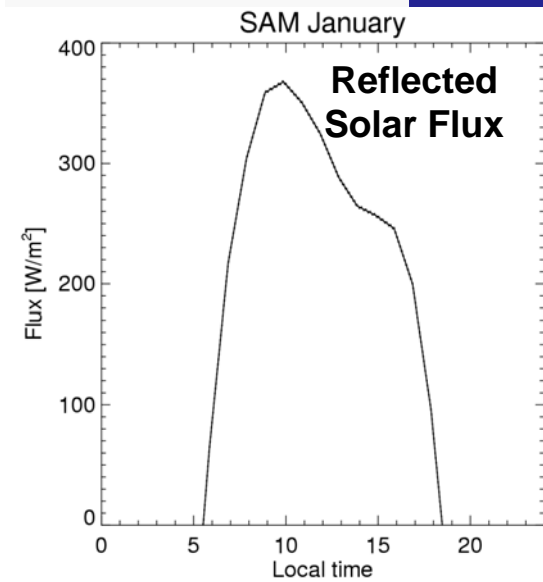
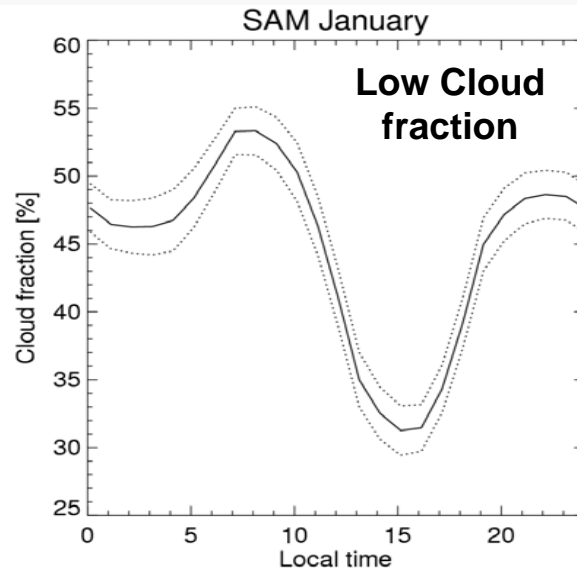
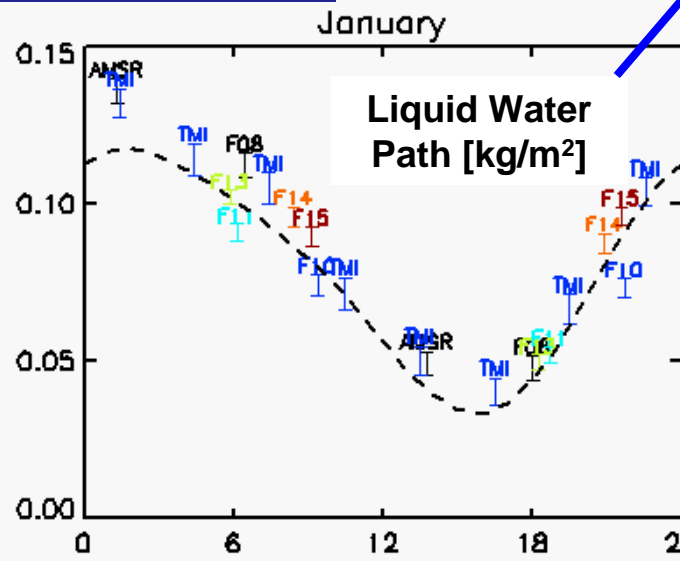
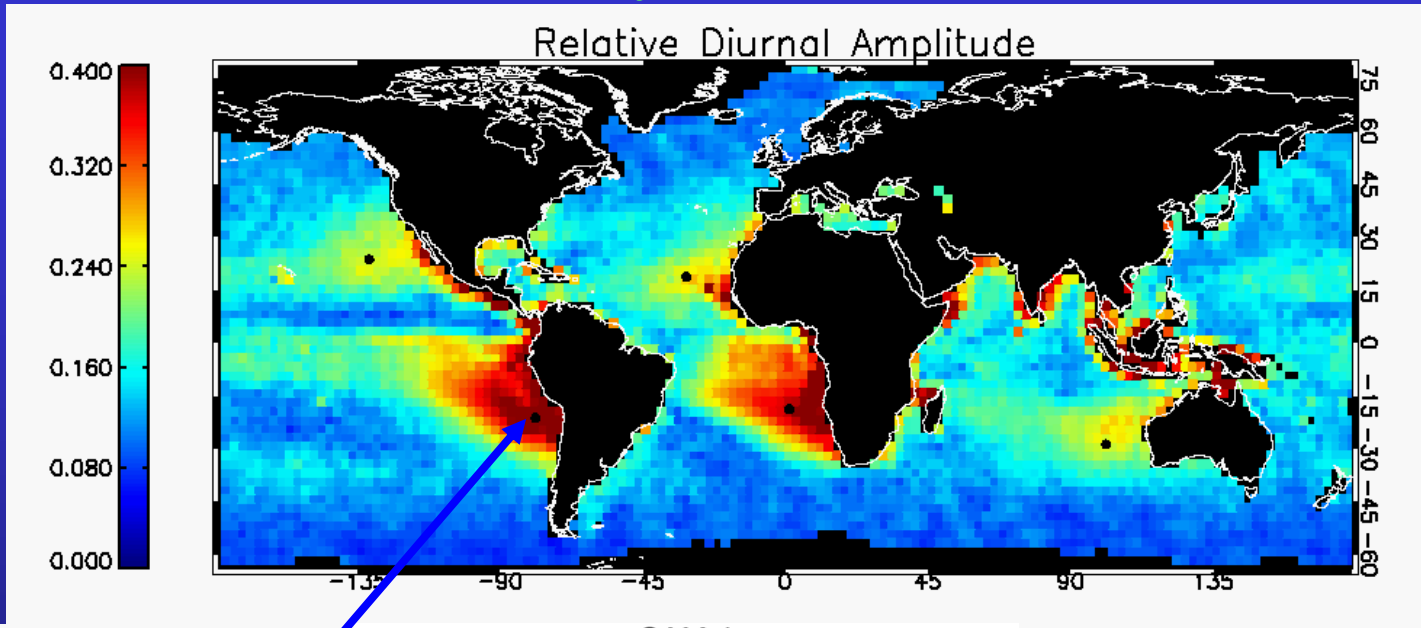


The diurnal cycle of clouds



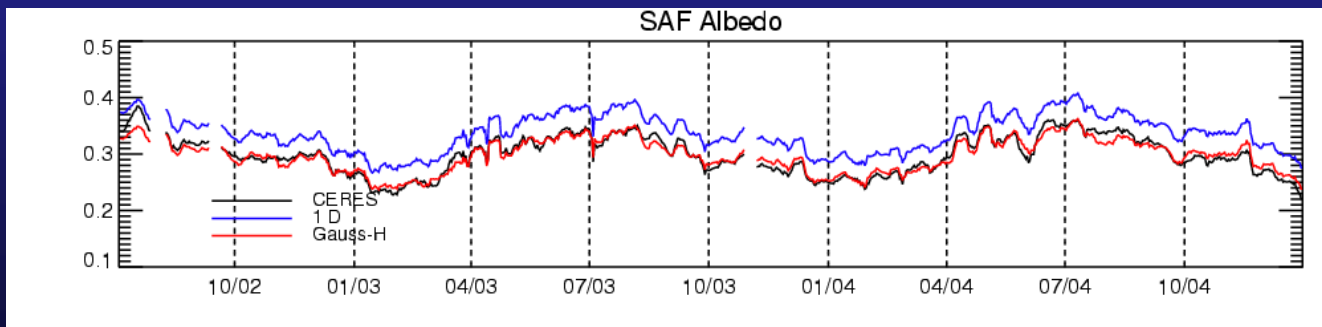
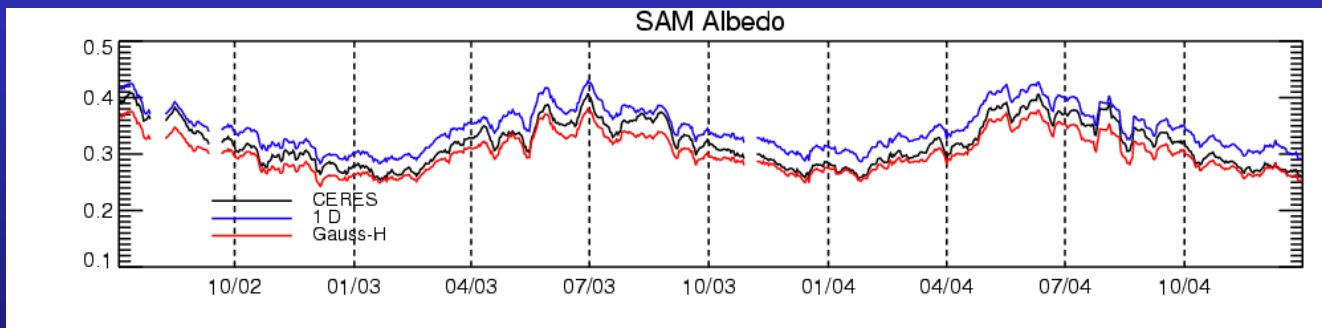
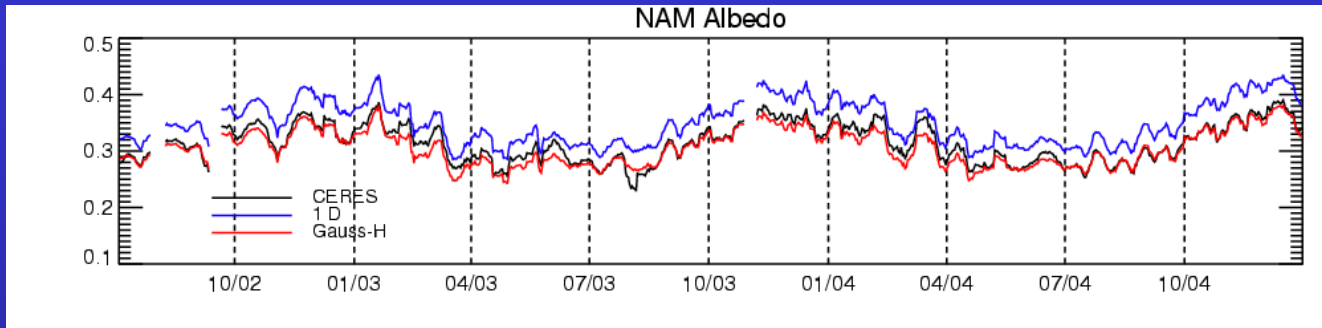
Local Time [hours]

The diurnal cycle of clouds

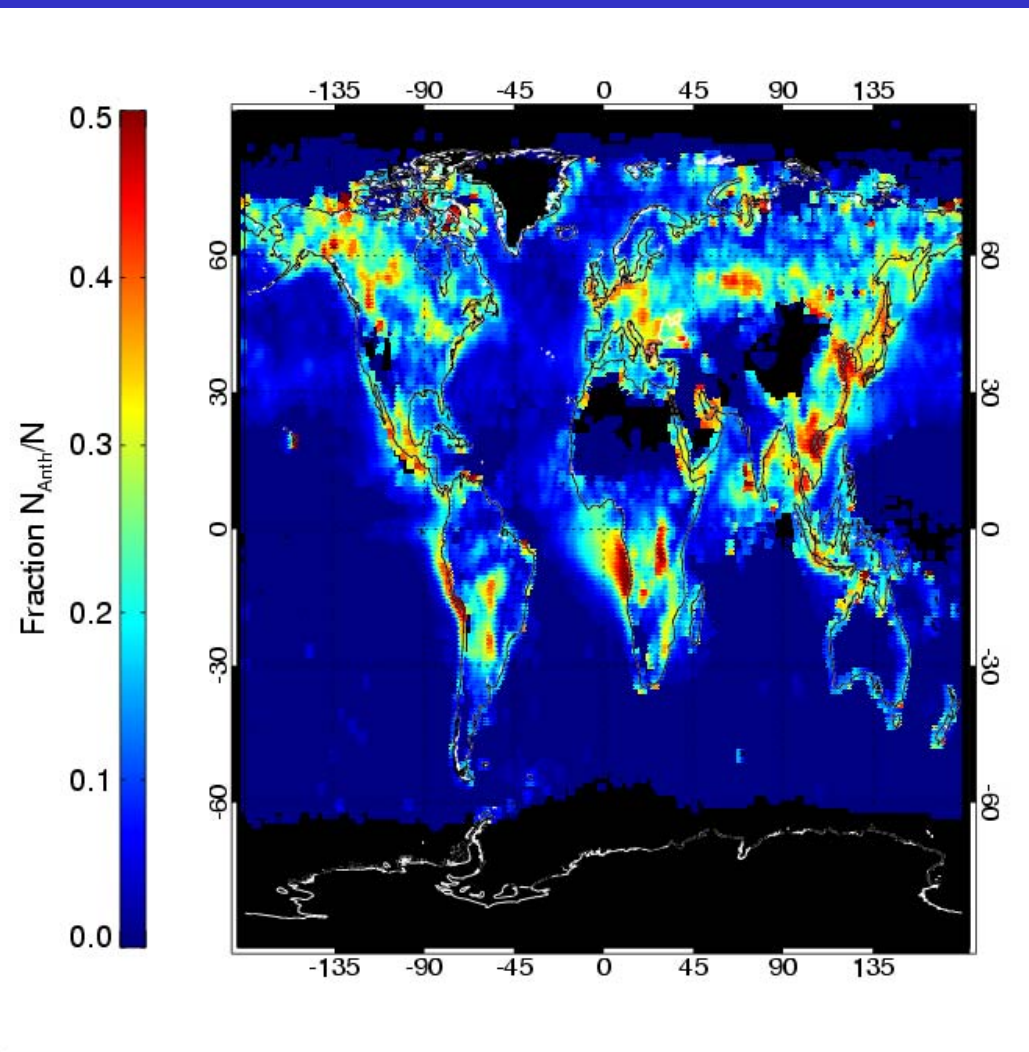


Local Time [hours]

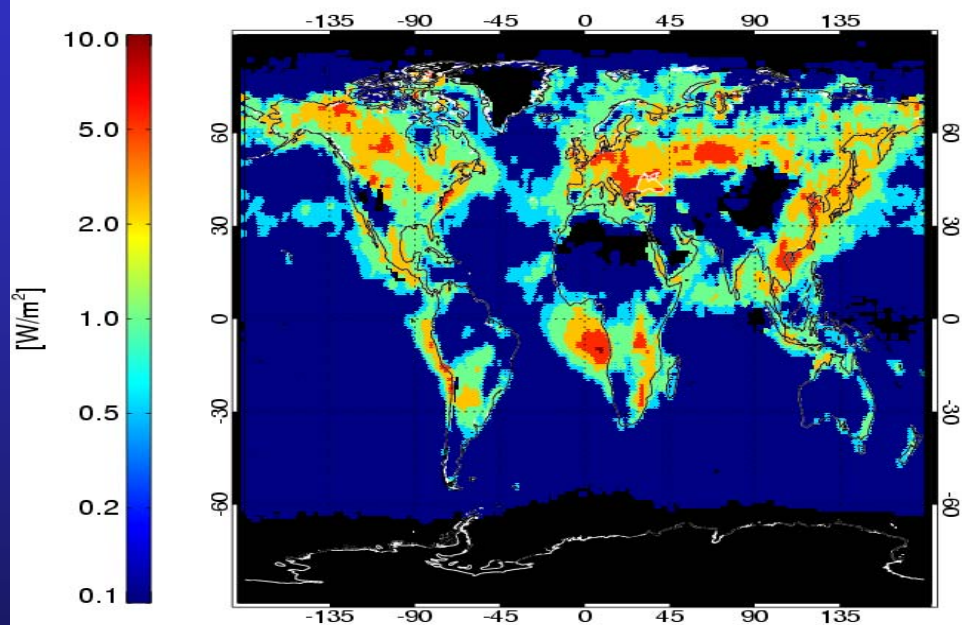
How does our albedo compare to observations (CERES)?



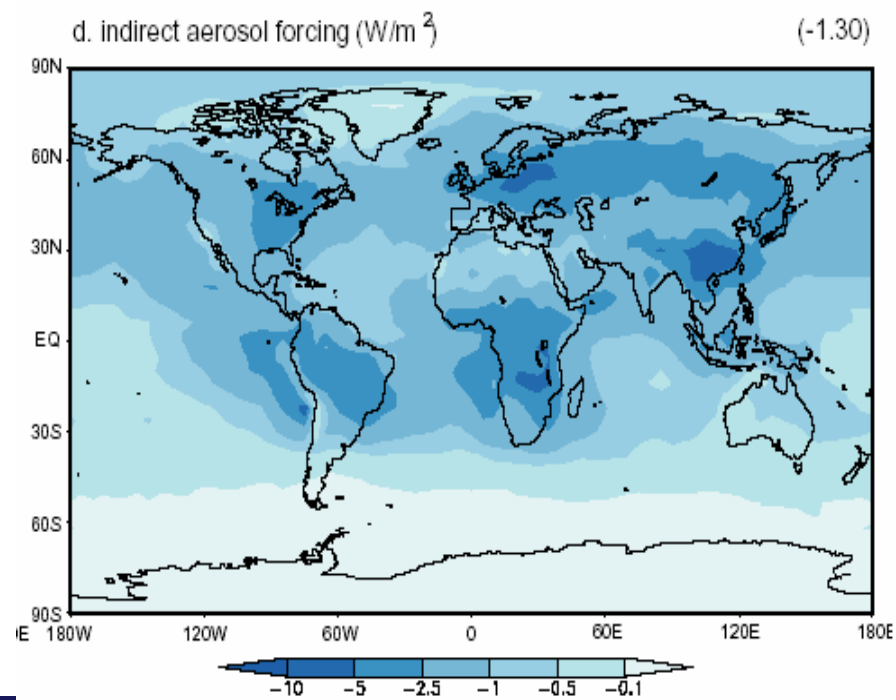
Anthropogenic fraction: Note: This is a **proxy**



IAE



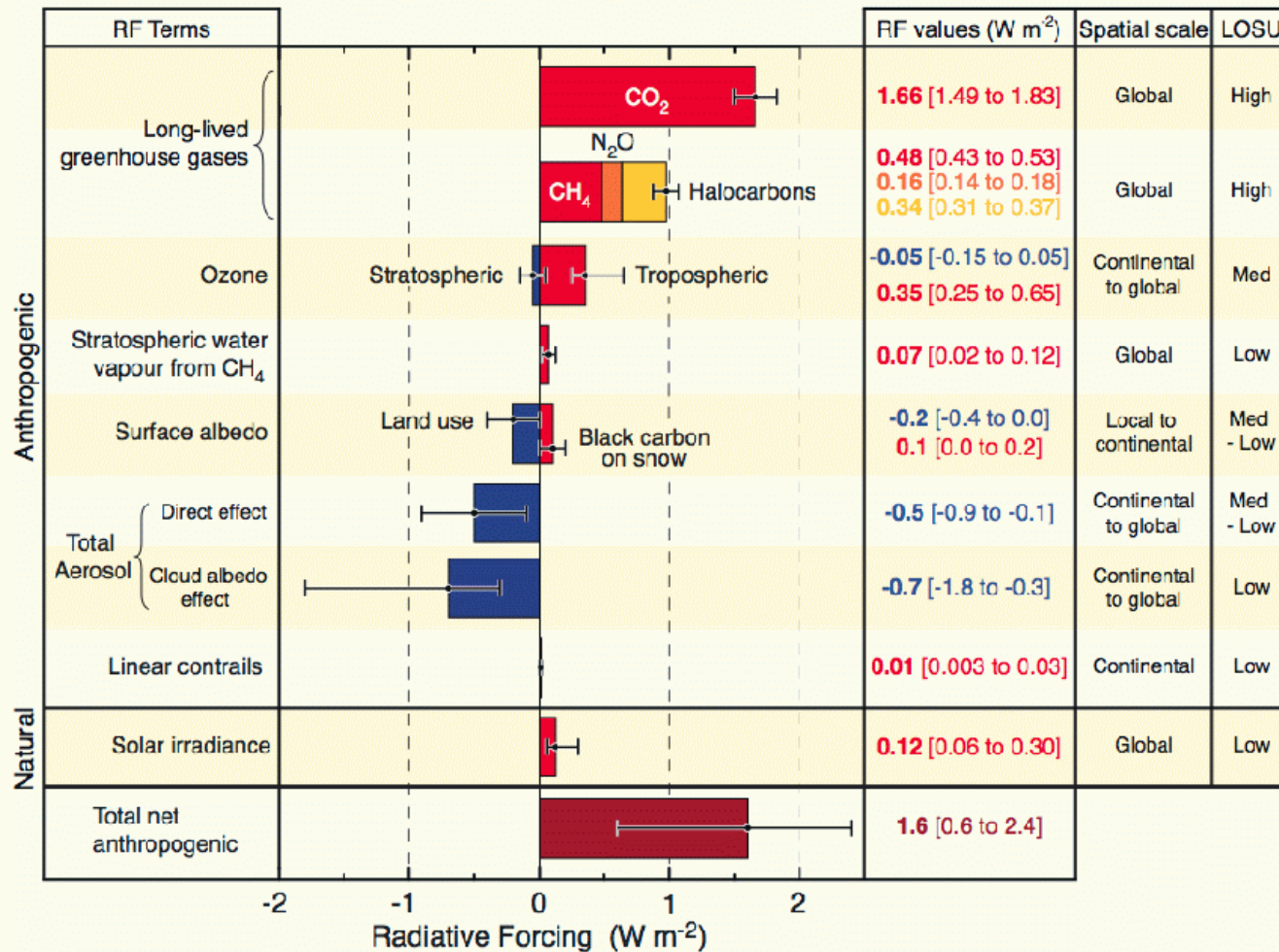
This study



(Chen and Penner, 2005, ACP)

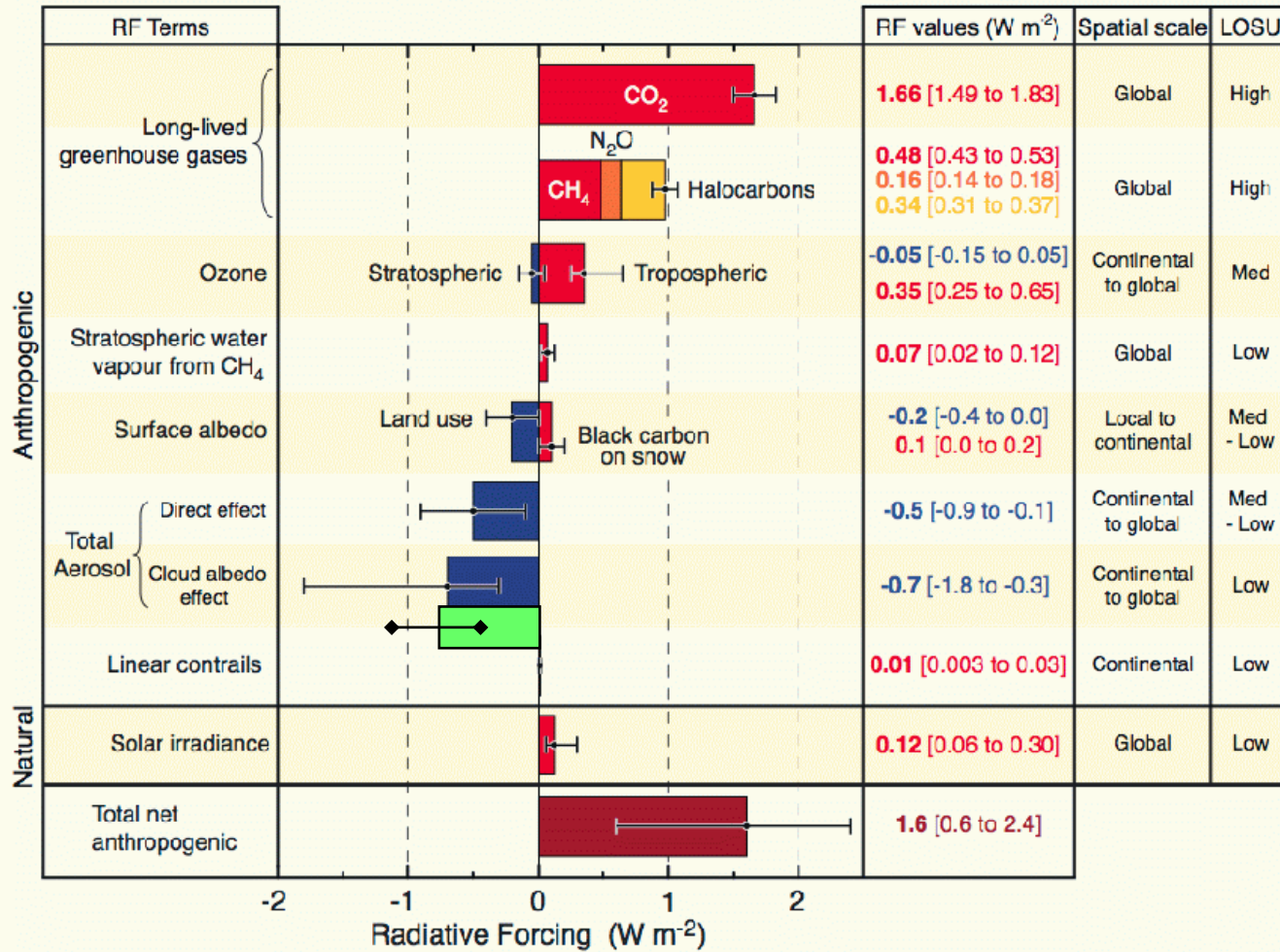
Conclusions

Radiative Forcing Components



Conclusions

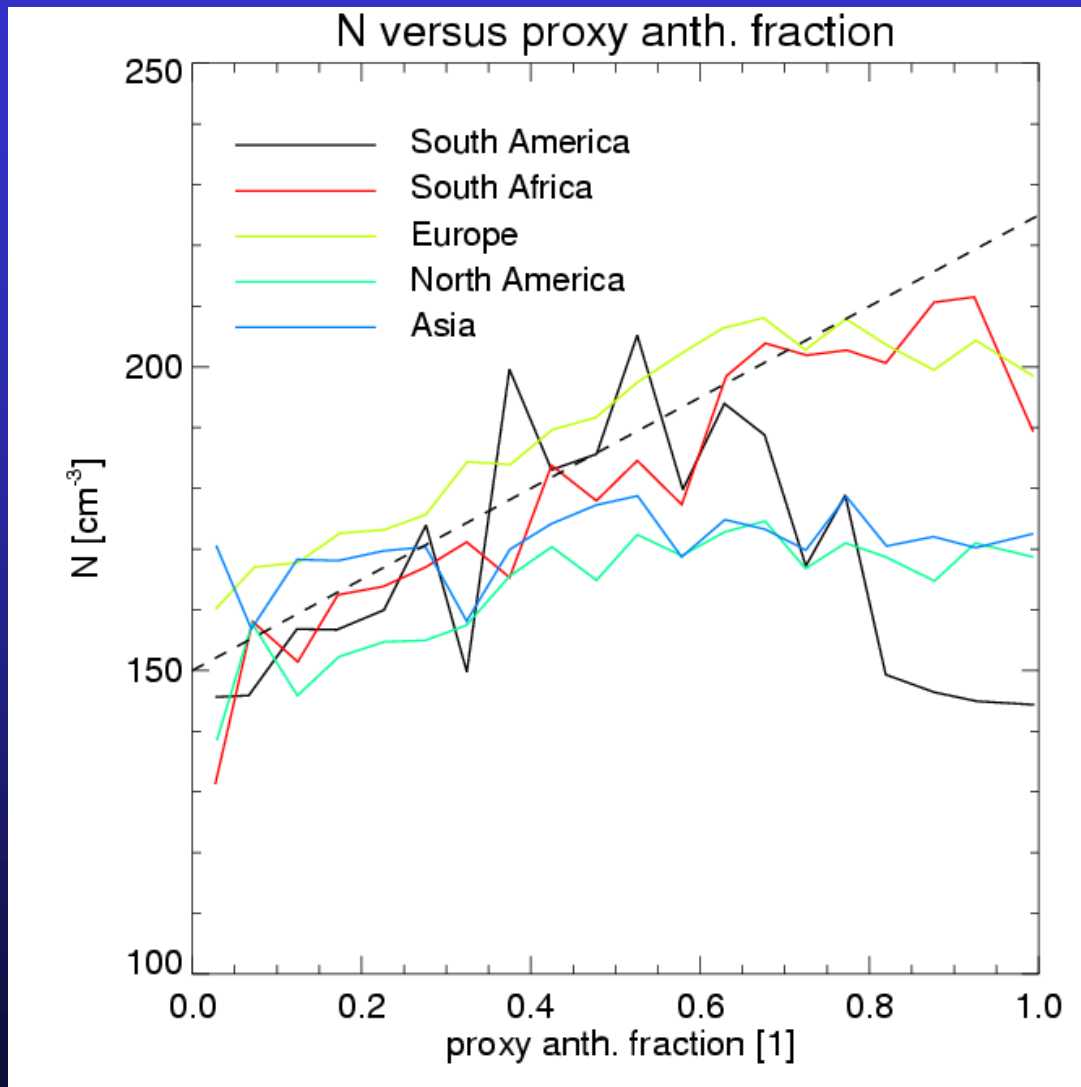
Radiative Forcing Components



Land

N to first order a function of anthropogenic aerosol optical thickness

Different behavior for different continents at high anthropogenic fraction



Ocean

$$N = 20 + 35 \cdot \left[\left(\log_{10} \left(c_{chlorophyll} \right) + 2 \right) \geq 0 \right] + 104 \cdot \frac{\tau_{anth,proxy}}{\tau}$$

